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THERMOANALYTICAL STUDY OF THE METAMORPHISM GRADE IN MALÉ KARPATY MTS. REGION

(Figs. 6)



Abstract: The paper deals with the problem of temperature and pressure conditions of the metamorphism in the productive zones of Pezinok-Pernek crystalline complex and in the black slates of Harmónia Group in Malé Karpaty Mts. using the results of thermal analysis of the rocks and the extracted organic matter (DTA in the controlled atmosphere of oxygen, nitrogen and static air). In conclusion the possibility in differentiation of these series is suggested.

Резюме: В статье решена проблема условий температуры и давления метаморфического процесса в продуктивных зонах Пезинок-Пернецкого кристалликума и в черных сланцах гармонской серии в Малых Карпатах на основе результатов термического анализа пород и химически отделенного органического вещества (измерения были проведены в контролируемой атмосфере O_2 , инертной атмосфере N_2 и в атмосфере статического воздуха). Внимание обращено на возможность дифференциации обеих толщ.

The geological structure of Pezinok-Pernek crystalline complex

Pezinok-Pernek crystalline complex represents bulky (4—8 km broad) zone of crystalline schists extending in NW-SE direction oriented approximately perpendicular to the range of mountains (C a m b e l, 1958). The specific conditions of the geological evolution in geosynclinal region of Lower Paleozoic formation in this series were described by C a m b e l (1958). This magmatogenic-sedimentary complex is bordered on its NE and SW side with granitoid rocks of Bratislava and Modra massifs (Fig. 1).

The part of Pezinok-Pernek crystalline complex consists of so called productive zones defined by C a m b e l (1959), (Fig. 1).

To the opinion of C a m b e l (1959) in these series are situated all antimonite deposits which are in essence analogous to the positions of synsedimentary submarine-volcanic-pyrrhotite lenses of the paleozoic age (C a m b e l, 1956; P o l á k, 1956).

The crystalline schists in the productive zones of Pezinok-Pernek crystalline complex consist of metamorphosed derivatives formed by combined metamorphism: at first by progressive periplutonic-regional metamorphism in Variscan orogeny, or by means of the contact metamorphism connected with the intrusion of granitoid magma followed by retrograde tectonometamorphic younger Alpine-Carpathian orogeny affecting not uniformly the schists of Malé Karpaty crystalline complex (Beck — V e t t e r s, 1904; K o u t e k — Z o u b e k, 1936; C a m b e l, 1962).

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The slaty strata group in Pezinok-Pernek crystalline complex occurs in its NE edge in the subgrade of the independent lithological-facial, stratigraphic younger series of Harmónia Group (Low Devonian — Carboniferous) (Cambel — Čorná, 1974), defined between Modra and Častá by Cambel (1954).

The exploration by boring has revealed that series of Pezinok-Pernek crystalline unit overlies the younger rocks of the Harmónia Group. It cannot be

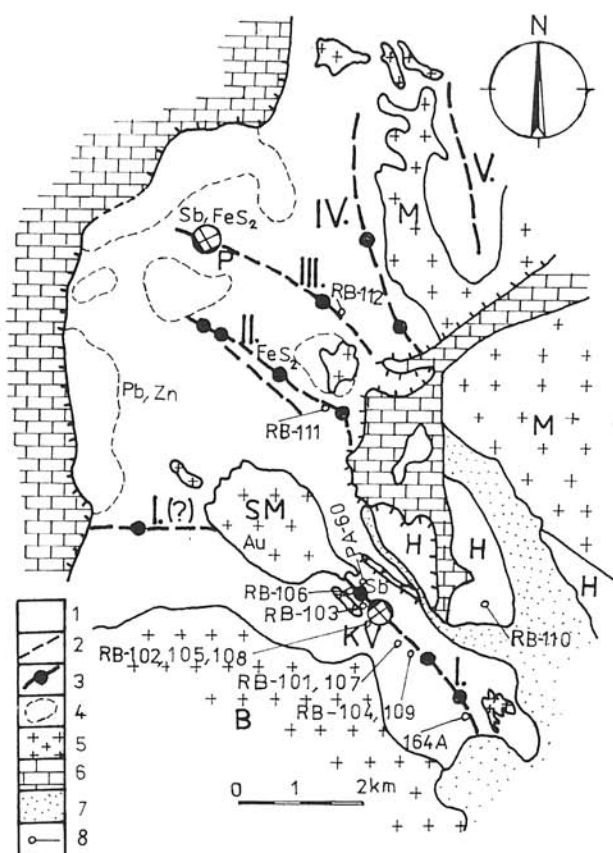


Fig. 1. Schematic geological map of the Pezinok-Pernek crystalline complex in Malé Karpaty region (Polák — Rak, 1980), inclusive the localization of the samples.

Explanations: 1 — metamorphosed rocks on the whole; 2 — productive zones (I—V); 3 — the occurrence of the Sb mineralization in productive zones; 4 — the regions with the insufficiently defined course of the productive zones; 5 — bigger bodies of granitoid rocks; 6 — Mesozoic rocks of the covering and higher tectonic units; 7 — Neogene rocks — Quaternary; 8 — localization and indicating of the sample; KV — Sb-deposit Pezinok-Kolársky vrch; P — Sb-deposit Pernek-Jahodnisko; B — Bratislava granitoid massif; M — Modra granitoid massif; SM — granitoid massif Staré mesto; H — Harmónia Group; Au, Pb, Zn, FeS — the occurrence of the corresponding mineralization.

decided whether a thrust fault is involved or whether the crystalline complex is allochthonous as is presumed by M a h e l' (1980, 1983).

The important part in Pezinok-Pernek and Harmónia Groups are the black slates. According to C a m b e l et al. (1980, 1981), C a m b e l — K h u n (1979), K h u n (1980) the black slates in productive zones of Pezinok-Pernek crystalline complex contain syngenetic (syndimentary) accumulations of the sulphide ores (pyrite-pyrrhotite). Antimonite ores represent accumulation which were mobilized during metamorphic and hydrothermal processes and then epigenetically deposited. The black slates were in Alpine-Carpathian orogeny partly affected by remobilization again and they were locally tectonically overworked. The black slates of Harmónia Group are not ore bearing.

The pressure and temperature conditions in the origin of the metamorphic rocks in Malé Karpaty region (concise review)

The first data on the pressure and temperature conditions (p, T conditions) in the origin of the metamorphic rocks in Malé Karpaty region were given by D y d a (1980 a, b; 1981, 1982) and C a m b e l et al. (1981, 1983). These authors have presumed for regional-periplutonic metamorphic process the temperature interval 450—610 °C and pressure 3.0—3.5 kbar.

K o r i k o v s k i j et al. (1984) in the study of mineralogical paragenesis of metamorphism in Malé Karpaty region distinguished two stages:

1. The regional-periplutonic stage defined by temperatures 500—550 °C and pressures 300—350 MPa reaching maximum (the highest P—T values) during intrusion of Bratislava granitoid magma. K o r i k o v s k i j et al. (1984) have divided the metamorphosed rocks in the crystalline complex of Malé Karpaty into four zones: biotitic, garnetic, staurolite-chloritic and staurolite-sillimanitic.

2. Metamorphism connected with the crystallization of Modra massif is characterized by pressure conditions 100—150 MPa and under its influence andalusite and cordierite cherts and spotted shales are formed.

Localization of the samples analyzed

The samples of the black slates represent first of all the productive zone in which the sole exploited deposit of the antimonite ores in Malé Karpaty region (Pezinok) is situated (Fig. 1). The prevailing part of the samples was collected from the geological bores performed by the group directed by S. P o l á k (Geological Survey, Spišská Nová Ves). The samples collected from the workings (antimonite gallery, PA—60) and exposures (RB—110, 164 A) were studied to compare the results.

Thermoanalytical measurements

Powder samples (below 0.05 mm) were studied using Dupont 990 Thermo-analyzer by means of DTA method (the sample weight: 15—20 mg, heating rate 20 °C min⁻¹) in dynamic oxygen and inert (N₂) atmosphere (the rate of gas flow was 1 cm³ s⁻¹).

The starting temperature "T₀" (the initiation of the oxidation of the organic matter) and peak temperature were derived from DTA curves obtained in

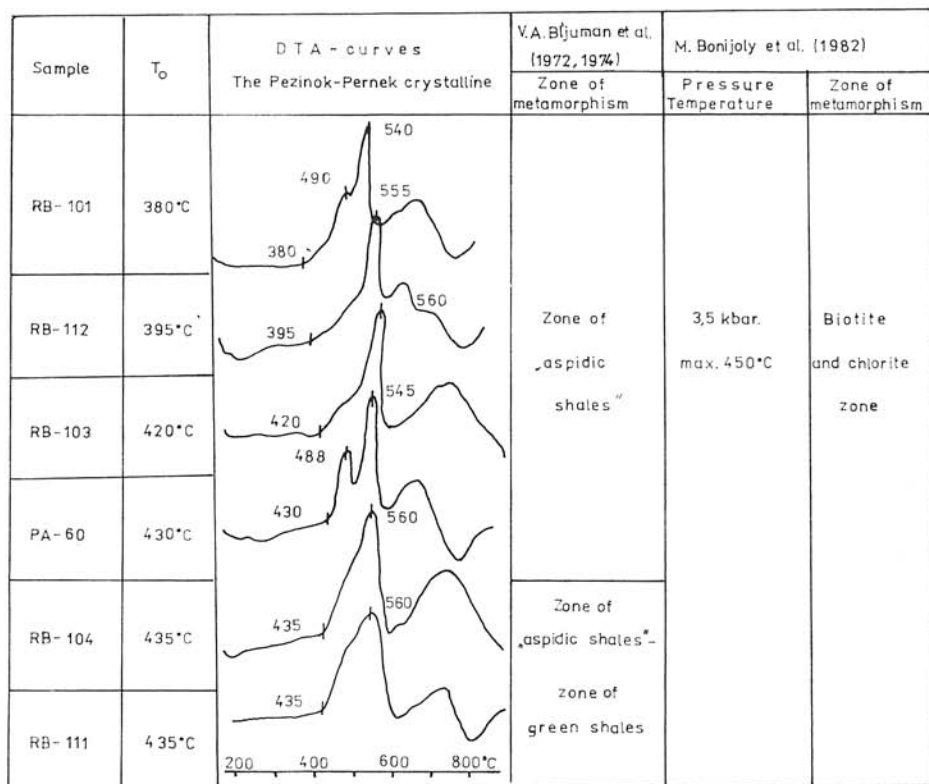


Fig. 2. Thermograms for the black states of Pezinok-Pernek crystalline complex, T_0 — the initiation of the oxidation reaction.

oxygen atmosphere (Figs. 2, 3). The same data were extracted from endothermic and exothermic effects from DTA curves obtained in nitrogen atmosphere (Fig. 5, 6).

Organic matter extracted from the black slates by means of the flotation in toluene according to the method by Bonijoly et al. (1982) was analyzed in the static air using Derivatograph MOM Q-1500 D (the sample weight: 0.5–1.0 g, heating rate 10°C min⁻¹). The results are shown in Fig. 4. The initiation of the exothermal peak occurs at temperatures 360–435°C, peak temperature occurs at 468–485°C.

Determination of the metamorphism grade in productive zones of Pezinok-Pernek crystalline complex

Bljumen et al. (1972, 1974) have characterized various metamorphic rocks containing the admixture of organic matter by means of the starting temperature of oxidation reaction using the thermoanalytical method. This temperature increases with increasing of metamorphosis and progressive graphi-

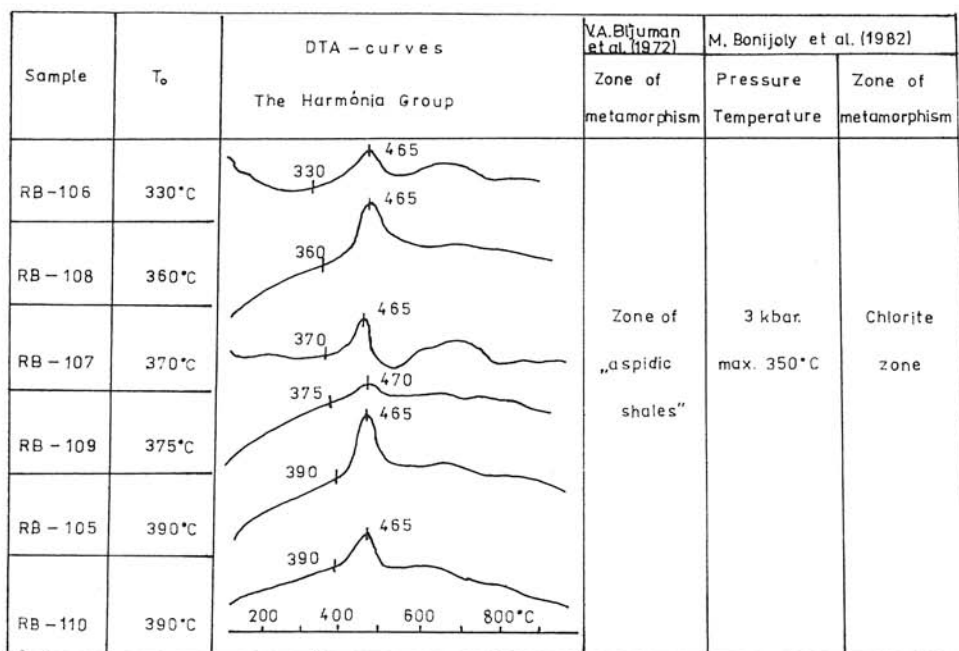


Fig. 3. Thermograms for the black states of Harmónia Group, T₀ — as in Fig. 2.

tization of carbonaceous matter in the rock. Bljumen et al., (1972, 1974) stated that the starting temperature of oxidation reaction does not depend upon the initial composition of the rock as well as upon amount of organic matter. The initial temperature of oxidation is about 100 °C higher for every next facia. Bljumen et al., have also found that the graphite structure formed at higher temperature does not change in heating at lower temperature.

The initiation of the first exothermal peak for the samples of Pezinok-Pernek series occurs in temperature interval 380—435 °C corresponding to the phase of the low metamorphosed “aspide-slates” (Bljumen et al., 1972, 1974) beyond the boundary of the green slates in the biotite-chlorite metamorphic zone with pressure conditions up to 3.5 kbar and maximum temperature up to 450 °C (Bonijoly et al., 1982). It concerns the rocks of low grade metamorphosis according to Winkler's classification (Winkler, 1974). Exothermal effect proceeds in the narrow temperature interval and exhibits a sharp maximum. The oxidation reaction finishes at the temperature lower than 900 °C which is typical for the paleozoic rocks as described by Bljumen et al. (1972).

DTA curves of the black slates of Harmónia Group show the initiation of the first exothermic peak at temperatures 330—390 °C what indicates likewise the black slates of Pezinok-Pernek Group the stage of low metamorphosed “aspide-slates”. The pressure and temperature conditions were, according to calculations and comparisons of Bonijoly et al. (1982), slightly lower (less than 3 kbar) therefore the slates of Harmónia Group may be classified

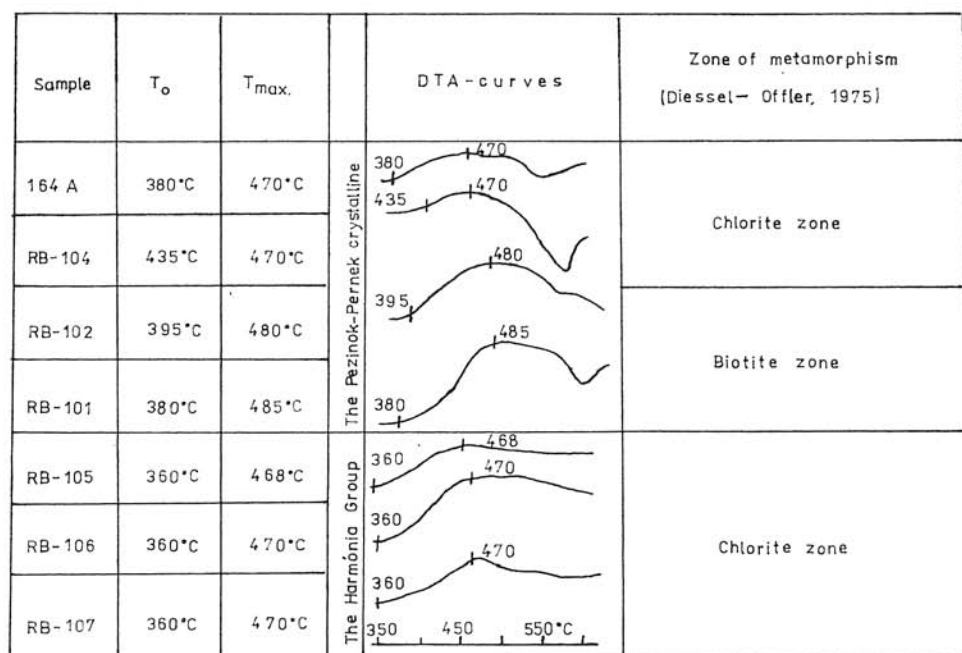


Fig. 4. Thermograms for the organic matter separated from the black states of Pezinok-Pernek and Harmónia Groups.
Explanations: T_0 — as in Fig. 2; $T_{max.}$ — DTA peak temperature.

as the slates of chlorite zone. According to Winkler (1974) and to the thermoanalytical data represent the slates of Harmónia Group the rocks of very low and low grade metamorphism. They can be classified one grade lower than the slates of Pezinok-Pernek Group which were metamorphosed in the conditions of the biotite-chlorite zone.

The DTA curves of organic matter extracted from the black slates by means of toluene (Fig. 4) confirmed the results obtained from the samples of the rocks.

The samples of Pezinok-Pernek Group on the basis of the maximum of the first exothermal peak represent the rocks of chlorite and biotite metamorphic zone (Diesel — Offler, 1975). Characterization of the metamorphism grade according to the methods by Bljumen et al. (1972, 1974) and Bonijoly et al. (1982) is the same as in the case of the samples of the rocks.

*The possibility of differentiation of the slates
 from Pezinok-Pernek and Harmónia Groups on the basis
 of thermoanalytical results*

The importance of the differentiation of the black slates from Pezinok-Pernek Group comprising a part of the productive zones form the Harmónia

Group has been suggested by Cambel — Khun (1979), Khun (1980), Cambel — Khun (1983) from the lithological point of view as well as from the standpoint of the genesis of the mineralization.

The samples of Pezinok-Pernek Group, analyzed in oxygen atmosphere, exhibit the exothermal peak in the temperature region 540—560 °C (Fig. 2). Less developed exothermal effect (with the exception of the samples PA-60 and RB-101) occurs at temperatures 470—490 °C. Intensive exothermal peak at 540—560 °C is typical for Pezinok-Pernek Group.

In temperature interval 600—800 °C occurs next exotherm. The initial and peak temperatures of this effect do change in the broad temperature interval.

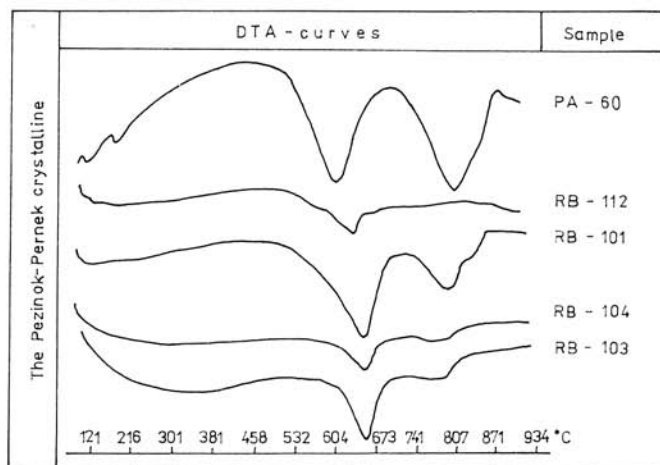


Fig. 5. Thermograms for the black slates of Pezinok-Pernek Group (DTA in inert atmosphere — N₂).

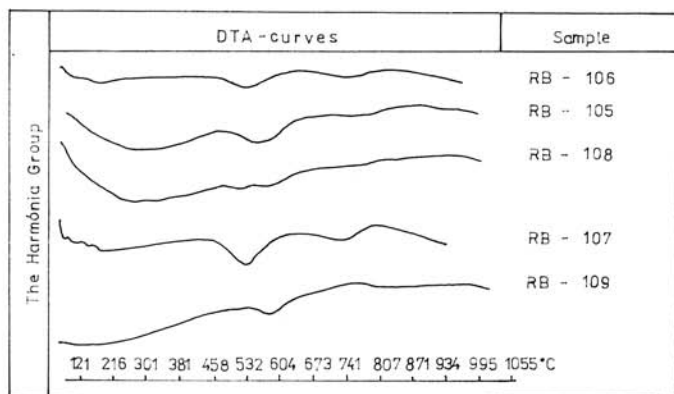


Fig. 6. Thermograms for the black slates of Harmónia Group (DTA in inert atmosphere — N₂).

The occurrence of this exotherm is also typical for all samples of Pezinok-Pernek Group. Endothermal effect at 800 °C which can be observed in DTA curves of the samples PA-60 and RB-101 represent the heat consumption during decarbonization reaction.

DTA curves obtained from the samples of Harmónia Group distinguish from the former significantly (Fig. 3). They are characteristic by lower intensity of the exothermal effects and do not exhibit maximum at 540—550 °C. The common for these curves is exotherm at 465—470 °C and less developed maximum in the region of 700 °C. The starting temperature of the first exothermal reaction occurs at 330—390 °C (Fig. 3).

Thermograms of the samples from Pezinok-Pernek Group analyzed in inert nitrogen atmosphere show two significant endotherms (in the exception of the sample RB-102) at 600—650 °C and 800 °C. The first endotherm indicates the process of the evolution of organic matter by distillation, the second one the decarbonization (Fig. 5).

The samples of Harmónia Group analyzed in nitrogen atmosphere distinguish from the former by morphology and temperature characteristics of the effects in DTA curves (Fig. 6).

Conclusion

On the basis of graphite geothermometer was studied the grade of metamorphism for the rocks of productive zones in Pezinok-Pernek crystalline complex and Harmónia Group by means of thermal analysis.

According to the results of thermal analysis the rocks of Pezinok-Pernek Group were metamorphosed in the stage of biotite-chlorite zone corresponding to the pressure 3.5 kbar and the maximum temperature 450 °C (Bonijoly et al., 1982). According to the classification of Winkler (1974) it concerns the rocks with low stage of metamorphosis. The data of thermal analysis show, that the highest temperature of the metamorphosis was 435 °C.

These results would evidence that the rocks of Harmónia Group were metamorphosed in the conditions of the chlorite zone at the temperatures 330—390 °C. The schists of the Harmónia Group belong to the slates on the boundary between very low and low grade metamorphosis according to the classification of Winkler (1974) and to the results of the thermal analysis.

The temperatures and pressures determined by the method of graphite thermometer are lower compared with data given by Dyda (1980 a, b, 1981, 1982), Cambel et al. (1981) and Korikovskij (1984).

The highest temperatures and pressures stated the papers by Dyda (1980 a, b; 1981; 1982) and Cambel et al. (1981). These works are based on the study of index minerals (they have determined the temperature interval 450—610 °C and pressures 300—350 MPa). Korikovskij et al. (1984) studying the mineral paragenesis have found lower parameters (500—550 °C, 100—350 MPa).

The thermoanalytical study of the black slates showed, that the lowest *p*, *T* conditions are given by the method of graphite geothermometer (330—435 °C, 300—350 MPa) when comparing with other methods.

This fact must be taken into consideration in the thermobarometric studies. Diversity of methods used in determination of p, T conditions of metamorphism of slates in the Malé Karpaty Mts. enabled such comparison.

Translated by V. Figuš

REFERENCES

- BECK, H. — VETTERS, H., 1904: Zur Geologie der Kleinen Karpathen, eine stratigraphisch — tektonische Studie. Beitr. Paläont. Geol. Österr., — Ung. Orients 16.
- BONIJOLY, M. — OBERLIN, M. — OBERLIN, A., 1982: A possible mechanism for natural graphite formation. Int. J. Coal geol. Elsevier Scientific Pub. Comp. (Amsterdam), 1, pp. 238—312.
- CAMBEL, B., 1954: K otázke kryštalickej bridlice medzi Cajlou a Hornými Orešami v Malých Karpatoch. Geol. Práce, Zpr. 1 (Bratislava), pp. 16—20.
- CAMBEL, B., 1956: Genetické problémy zrudnenia v Malých Karpatoch. Geol. Práce, Zpr. 9 (Bratislava), pp. 5—27.
- CAMBEL, B., 1958: Príspevok ku geológii pezinsko — perneckého kryštalinika. Acta geol. geogr. Univ. Comen., Geol. 1, pp. 137—260.
- CAMBEL, B., 1959: Hydrotermálne ložiská v Malých Karpatoch, mineralógia a geochemia ich rúd. Acta geol. geogr. Univ. Comen., Geol. 3, pp. 338.
- CAMBEL, B., 1962: In: Buday, T. — Cambel, B. — Maheľ, M. et al.: Vysvetlivky k prehľadnej geologickej mape ČSSR 1 : 200 000, M — 33 — XXXV, M — 33 — XXXVI, Wien — Bratislava. Geofond 1962 vo Vyd. SAV, pp. 248.
- CAMBEL, B. — KHUN, M., 1979: Distribúcia a korelácia stopových prvkov v čiernych bridliciach kryštalinika Malých Karpát. Miner. slov. (Bratislava), 11, 6, pp. 507—520.
- CAMBEL, B. — KÁTLOVSKÝ, V. — KHUN, M., 1981: Geochemia uránu, tória, uhlíka a ďalších prvkov v tmavých bridliciach kryštalinika Malých Karpát. Miner. slov. (Bratislava), 15, 5, pp. 423—441.
- CAMBEL, B. — DYDA, M. — SPIŠIAK, J., 1981: Thermodynamic measurements of minerals in metamorphites in the area of crystalline of Malé Karpaty Mts. Geol. Zbor. Geol. carpath. (Bratislava), 32, 6, pp. 745—760.
- CAMBEL, B. — KHUN, M., 1983: Geochemical characteristics of black shales from the ore-bearing complex of strata of the Malé Karpaty Mts. Geol. Zbor. Geol. carpath. (Bratislava), 34, 3, pp. 255—382.
- DIESSEL, C. F. K. — OFFLER, R., 1975: Change in physical properties of coalified and graphitised phytoclasts with grade of metamorphism. Neu. Jb. Mineral., Abh.; Mh., 1 (Stuttgart), pp. 11—26.
- DYDA, M., 1980 a: Physical properties and temperatures of crystallization of coexisting garnets and biotites from parageneses of the Little Carpathians. Geol. Zborn. Geol. carpath. (Bratislava), 31, 1—2, pp. 201—213.
- DYDA, M., 1980 b: Metamorphic grade and packing index in garnets. Geol. Zbor. Geol. carpath. (Bratislava), 31, 3, pp. 359—374.
- DYDA, M., 1981: Metamorphic grade and packing index of coexisting biotites and garnets from the Malé Karpaty Mts. metapelitic rocks. Geol. Zbor. Geol. carpath. (Bratislava), 32, 2, pp. 269—286.
- DYDA, M., 1982: Koexistencia granátu a biotitu v malokarpatských metapelitoch. Zborník „Metamorfne procesy v Západných Karpatoch“ GÜDS, Bratislava, pp. 69—73.
- KHUN, M., 1980: Použitie diskriminačnej analýzy na príklade čiernych bridlic malokarpatského kryštalinika. Miner. slov. (Bratislava), 12, 5, pp. 469—473.
- KOUTEK, J. — ZOUBEK, V., 1936: Vysvetlivky ke geologické mape v měřítku 1 : 75 000, list Bratislava. Knihovna SGÚ, Praha.
- MAHEĽ, M., 1980: Sú granitoidné masivy Malých Karpát príkrovmi? Miner. slov. (Bratislava), 12, 2, pp. 185—187.
- MAHEĽ, M., 1983: Križňanský príkrov, príklad polyštruktúrnej jednotky. Miner. slov. (Bratislava), 15, 3, pp. 193—216.

- POLÁK, S., 1956: Niekoľko poznámok k otázke vzájomného vzťahu medzi pyritom a pyrotínom v malokarpatských kyzových zrudneniach. *Geol. Práce, Zpr.* 6 (Bratislava), pp. 41—44.
- WINKLER, H. G. F., 1974: *Petrogenesis of metamorphic rocks*. Springer Verl. New York — Heidelberg — Berlin, 327 pp.
- БЛЮМАН, В. А. — ДЬЯКОНОВ, И. С. — КРАСАВИНА, Т. Н., 1972: Изменение структурного состояния графита при прогрессивном региональном метаморфизме. Докл. Акад. Наук. СССР (Москва), 206, 5, pp. 1198—1200.
- БЛЮМАН, В. А. — ДЬЯКОНОВ, И. С. — КРАСАВИНА, Т. Н. — ПАВЛОВ, М. Г., 1974: Использование термо- и рентгенографических характеристик графита для определения уровня и типа метаморфизма. Зап. Всесоюз. минерал. Общ., Сер. 2 (Ленинград), 1, СП, pp. 95—103.
- КОРИКОВСКИЙ, С. П. — ЦАМБЕЛ, Б. — МИКЛОШ, Я. — ЯНАК, М., 1984: Метаморфизм кристалликума Малых Карпат: этапы, зональность, связь с гранитоидами. *Geol. Zbor. Geol. carpath. (Bratislava)*, 35, 4, pp. 437—462.
- ЦАМБЕЛ, Б. — ЖУКОВ, Ф. С. — САВЧЕНКО, Д., 1980: Генетические и изотопно-геологические особенности формирования кольчеданных руд в Малых Карпатах. *Miner. Slov. (Bratislava)*, 12, 6, pp. 540—553.
- ЦАМБЕЛ, Б. — ЧОРНА, О., 1974: Стратиграфия кристаллического основания Малых Карпат в свете палинологических исследований. *Geol. Zbor. Geol. carpath. (Bratislava)*, 25, 2, pp. 231—241.

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